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# **SMB from RCMs: structural uncertainties in sea level projections from both ice sheets**









NORCE







#### **Surface Mass balance of ice sheets: atmospheric and surface coupling**









Frozen lake surface



### **Atmosphere, snow and firn processes required for SMB:**





 $SMB = RF + SF - RU - SU + DE$ 

 $Runoff = Melt + Rainfall + Condensation - Retention - Refreeze$ 

 $SEB = SWD - SWU + LWD - LWU + SHF + LHF + GHF$ 



# **Present Day SMB from RCMs with Climate Reanalysis: GrIS**













#### Ensemble Mean

## **Present Day SMB from RCMs with Climate Reanalysis: AIS**

**DMI** 



sparse

![](_page_4_Picture_89.jpeg)

![](_page_4_Picture_3.jpeg)

# **Greenland Regional Climate Models:**

#### 15km resolution

#### Irreducible water saturation  $= 7\%$

Albedo scheme adjusts for snow properties and clouds constant in bare ice zone at 0.55.

> Irreducible water saturation  $= 2\%$

11km resolution

Snow albedo based on snow impurities, zenith angle, and metamorphism, bare ice albedo is derived from a MODIS 5% lowest values, averaged over 2000-2015.

HIRHAM5Albedo: linear ramping of snow albedo ranging from 0.85 below −5 °C to 0.65 at 0 °C, bare ice remains constant at 0.4 with transition albedo is determined for thin snow layers on ice.

> Irreducible water saturation  $= 7\%$

![](_page_5_Figure_15.jpeg)

#### 5km resolution

![](_page_5_Picture_200.jpeg)

# **MAR RACM02.3p2**

**A range of future projections are available..** 

# **CESM2 SSP5-8.5 as common forcing: GrIS**

![](_page_6_Figure_1.jpeg)

![](_page_6_Picture_2.jpeg)

![](_page_6_Figure_3.jpeg)

 $-250$ 

 $-1000$ 

 $-2000$   $\geq$ 

 $-3000$ 

 $-4000$ 

![](_page_6_Figure_4.jpeg)

2081- 2100

![](_page_6_Figure_6.jpeg)

![](_page_6_Figure_7.jpeg)

From Quentin Glaude

![](_page_6_Picture_9.jpeg)

## **Differences between models are driven by melt and runoff**

![](_page_7_Figure_1.jpeg)

![](_page_7_Figure_2.jpeg)

![](_page_7_Figure_3.jpeg)

RCMs have non-uniform response to temperature anomalies.

RACMO and HIRHAM are higher than MAR at low temperatures but MAR has steeper curve up at higher temperatures From Quentin Glaude

![](_page_7_Picture_11.jpeg)

#### RACMO: 82% melt to runoff

![](_page_7_Figure_5.jpeg)

![](_page_7_Picture_7.jpeg)

3500

3000

2500

 $\frac{1}{2}$  2000<br> $\frac{1}{2}$  1500<br> $\frac{1}{2}$  1500

1000

 $\mathbf{0}$ 

 $-2$ 

![](_page_7_Figure_8.jpeg)

Radiation and cloud parameterisations likely account for some differences:

- Mixed phase clouds (cloud cover and cloud optical depth)
- Surface albedo scheme differences

![](_page_8_Figure_4.jpeg)

![](_page_8_Picture_5.jpeg)

## **Surface Energy Budgets have compensating errors (GrIS)**

Liquid clouds reflect SWin and absorb LWout radiation

![](_page_8_Figure_9.jpeg)

Ice clouds reflect less SWin and absorb less LWout radiation

![](_page_8_Picture_11.jpeg)

![](_page_8_Picture_12.jpeg)

![](_page_8_Picture_13.jpeg)

From Quentin Glaude

![](_page_8_Figure_7.jpeg)

### **Different model sensitivities to temperature change**

![](_page_9_Figure_1.jpeg)

![](_page_9_Figure_2.jpeg)

**BIART** ERYOSPHERE & SEA LEVE

Melt – albedo feedback, but also IWS is key!

From Quentin From Quentin Glaude<br> **Glaude** 

![](_page_9_Picture_10.jpeg)

![](_page_9_Figure_11.jpeg)

![](_page_9_Picture_12.jpeg)

Ablation area has different sensitivities to temperature

RACMO: higher runoff from smaller fraction of area

![](_page_9_Picture_8.jpeg)

#### **Preprint out now:**

A Factor Two Difference in 21st-Century Greenland Ice Sheet Surface Mass Balance Projections from Three Regional Climate Models for a Strong Warming Scenario (SSP5-8.5)

Quentin Glaude<sup>1</sup>, Brice Noël<sup>2</sup>, Martin Olesen<sup>3</sup>, Michiel R. van den Broeke<sup>4</sup>, Willem Jan van de Berg<sup>5</sup>, Ruth Mottram<sup>3</sup>, Nicolaj Hansen<sup>3</sup>, Alison Delhasse<sup>1</sup>, Charles Amory<sup>6</sup>, Christoph Kittel<sup>7</sup>, Heiko Goelzer<sup>8</sup>, and Xavier Fettweis<sup>9</sup>

<sup>1</sup>Universite de Liege <sup>2</sup>Laboratoire de Climatologie et Topoclimatologie <sup>3</sup>Danish Meteorological Institute <sup>4</sup>Utrecht University <sup>5</sup>University of Utrecht  ${}^{6}$ Institut des Geosciences de l'Environnement <sup>7</sup>University of Liège <sup>8</sup>NORCE Norwegian Research Centre, Bjerknes Centre for Climate Research, Bergen, Norway <sup>9</sup>University of Liege, Belgium

![](_page_10_Picture_4.jpeg)

### **doi.org/10.22541/essoar.172537578.84386972/v**

![](_page_10_Picture_7.jpeg)

## **Diverging future projections over Antarctic ice shelves**

![](_page_11_Figure_1.jpeg)

**CRYOSPHERE & SEA LEVEL** 

![](_page_11_Picture_4.jpeg)

![](_page_12_Figure_1.jpeg)

![](_page_12_Picture_4.jpeg)

![](_page_12_Picture_5.jpeg)

## **Different model sensitivities to temperature change**

# **Surface melt and Cloud phases**

![](_page_13_Figure_1.jpeg)

From Hansen et al 2024

![](_page_13_Picture_3.jpeg)

![](_page_13_Figure_4.jpeg)

**180** 

 $180^\circ$ 

![](_page_13_Figure_6.jpeg)

![](_page_13_Figure_7.jpeg)

CERES 14th / CIWP 85.5

CERES 14th / CLWP

 $0.5$ 

![](_page_13_Figure_10.jpeg)

 $70°S$  $70^\circ S$   $\circ$ 

![](_page_13_Figure_12.jpeg)

CERES 17th / CIWP

0.01 0.02 0.05 0.1 0.2 0.5

#### MetUM 14th / CLWP

![](_page_13_Figure_18.jpeg)

#### MetUM 17th / CLWP

![](_page_13_Figure_20.jpeg)

![](_page_13_Picture_22.jpeg)

![](_page_14_Picture_5.jpeg)

![](_page_14_Picture_0.jpeg)

**Global climate models rarely produce SMB over ice sheets and regional climate simulations produce similar SMB values at present day Small differences between regional climate model physics parameterisatons can lead to large long-term large differences in SMB** 

SMB emulation and new model parameterisations under development for CMIP7 Summary SMB emulation and new model parameterisations under development for CMIP7 **We have a \*large\* amount of RCM projections over both Greenland and Antarctica forced by CMIP6 all publically available for analysis**

**projections, outside of driving climate induced uncertainty**

**Differences in radiation schemes and cloud parameterisations are important but s mall differences in firn and snowpack parameterisations can lead to long-term differences in retention, refreezing and runoff**