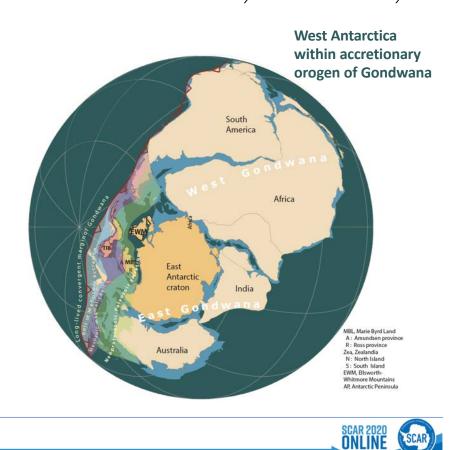
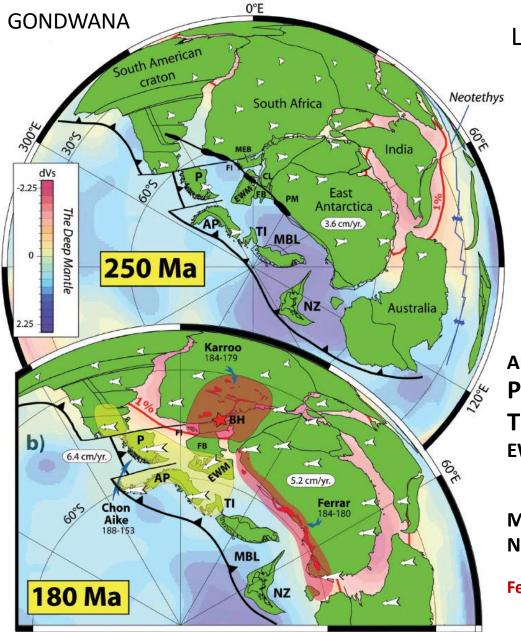


Updated tectonic framework of West Antarctica and legacy of formation within the complex convergent margin of Gondwana

> Christine Siddoway, Colorado College Tom A. Jordan & Teal Riley, BAS Kirsty Tinto, LDEO – Columbia University Matt Tankersley – Victoria University





Longstanding view of West Antarctica

Discrete tectonic entities: *Terranes* (character of intervening crust, not known)

Developed along the active margin of Gondwana

This presentation explores both distinctions and commonalities between three provinces of West Antarctica.

AP, Antarctic Peninsula

P, Patagonia

TI, Thurston Island

EWM, Ellsworth-Whitmore Mountains

> paleomagnetic evidence for large vertical axis rotation

MBL, Marie Byrd Land **NZ**, Zealandia

Ferrar – Large Igneous Province



Torsvik et al. 2008, doi:10.3133/of2007-1047.kp11

Presentation Outline

Part I.

I. Geological/geophysical/ geochronological evidence of a shared history for West Antarctica tectonic provinces along a complex accretionary margin

II. Geophysical remote sensing reveals the subglacial extent of geological provinces; new tectonic scenarios emerge.

- Ellsworth-Whitmore province (Jordan et al. 2017)
- Antarctic Peninsula (Burton-Johnson & Riley, 2015)
- Ross Embayment cratonic margin province (Tinto et al. 2019)

Part II.

III. The inboard limit of West Antarctica in Ross Embayment

- Defined by physiography or crustal geology?
- Transantarctic Mountains or Central High / boundary of cratonic margin province?

IV. Utility of updated crustal geology and tectonic framework Relevant to many areas of contemporary research: paleotopography, glacial processes, subglacial hydrology, heat flux, glacial isostatic response





SCAR members and programmes that contribute to this work



Geosciences Scientific Research Programmes

- Past Antarctic Ice Sheet Dynamics PAIS
- Solid Earth Responses & influences on Cryospheric Evolution-SERCE
- Antarctic Offshore Stratigraphy Project ANTOSTRAT

Geosciences Expert Groups

- Antarctic Digital Magnetic Anomaly Map Project –
- Antarctic Volcanism AntVolc
- Geodetic Infrastructure of Antarctica GIANT
- International Bathymetric Chart of the Southern Ocean IBCSO

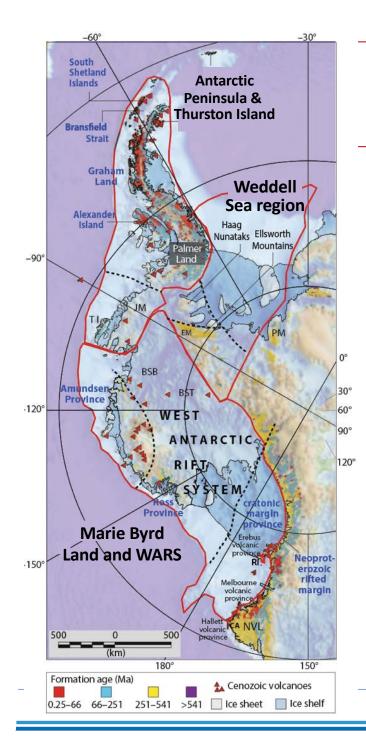
Geosciences Action Groups

- Connecting Geophysics with Geology CGG
- Geological Mapping Update of Antarctica –

ADMAP







Updated tectonic framework of West Antarctica and legacy of formation within the complex convergent margin of Gondwana

Weddell Sea region and continental highlands

(Ellsworth-Whitmore Mountains, Haag Nunataks) Structural trend of EWM is orthogonal to Pensacola and Transantarctic Mountains

Antarctic Peninsula (AP) – Thurston Island (TI) province

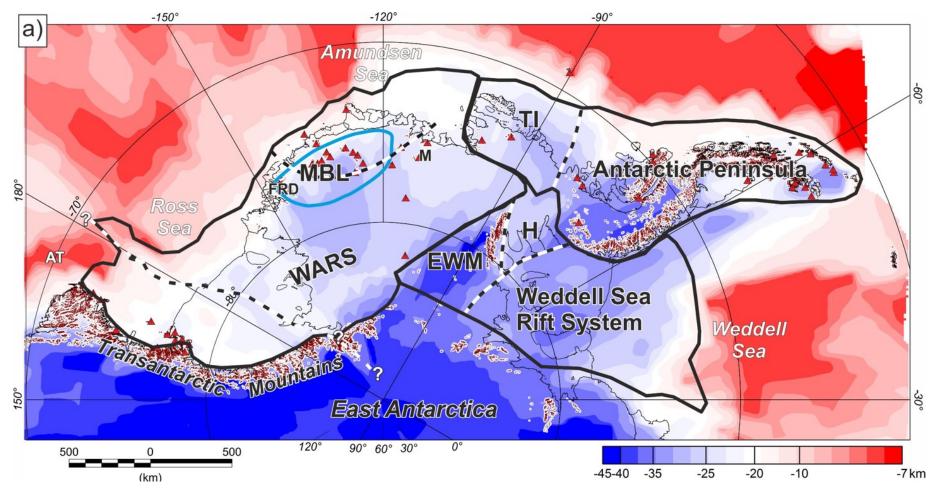
Marie Byrd Land (MBL) and West Antarctic rift system

Geophysical remote sensing provides new understanding of the subglacial extent lithospheric characteristics

locations ± kinematics of crustal structures

Geology, geochronology, and thermochronology has led to discovery of common elements shared among the three provinces.





Crustal thickness derived from passive seismic observations

West Antarctic crust has a mean thickness of just 25 km, in contrast to East Antarctica 's cratonic crust, 35–45-km-thick. The crustal thickness domains, together with aeromagnetic anomaly patterns, allow the subglacial extent of West Antarctic provinces to be determined. The geophysical characteristics aid the interpretation of the geological makeup of the crust.

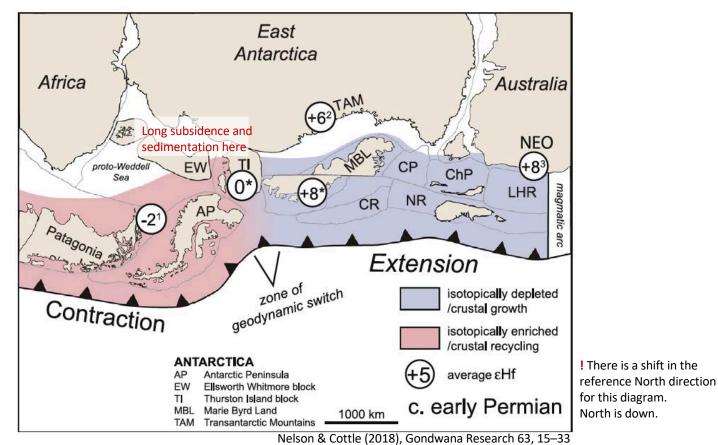
Thicker crust in the Peninsula - Weddell Sea region, compared to MBL, signifies contraction and tectonic thickening in the 'advancing' orogen.

Figure draws upon POLEnet contributions, including Chaput et al. 2014; An et al. 2015 and other POLEnet works



Gondwana accretionary orogen -

West Antarctica, Zealandia and Patagonia developed along this margin

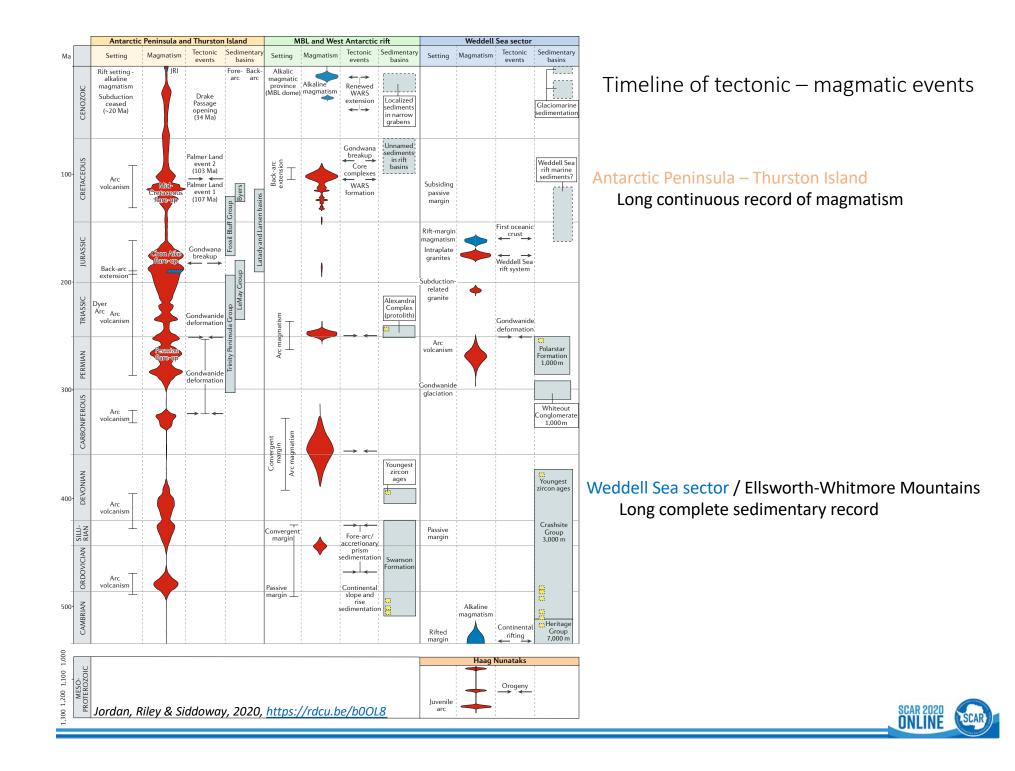


Accretionary orogens

- form at intraoceanic and continental margin convergent boundaries
- comprise supra-subduction zone forearc, magmatic arc and back-arc
- Retreating mode: marked by long-term extension, with development of back-arc basins. Subducting lower plate retreats with respect to overriding plate. Example: western Pacific, today.
- Advancing mode: marked by contraction and crustal thickening, with development of foreland fold and thrust belts. Overriding plate advances towards the downgoing plate (e.g. Andes)

See Cawood et al. 2009, GSL-SP 318

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Weddell Sea province

Generally long-wavelength magnetic features.

Northern Weddell Magnetic Province, offshore Southern Weddell Magnetic Province beneath Ronne Filchner ice shelf (RFIS)

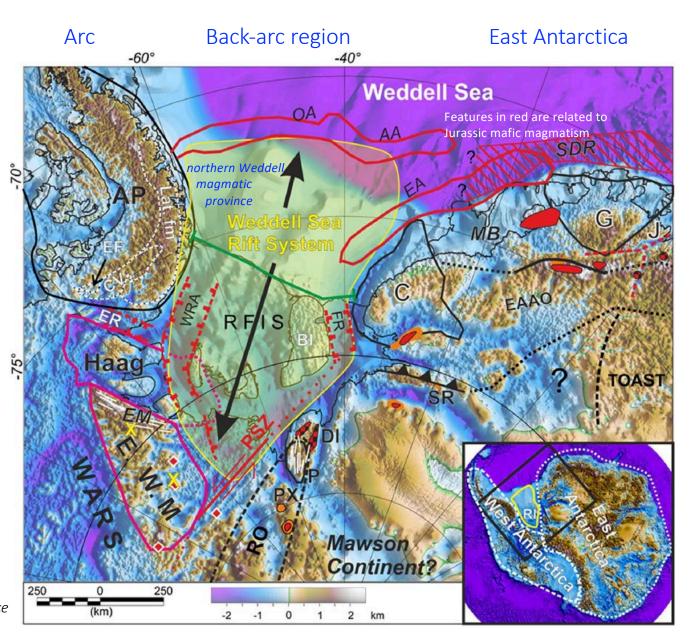
crust ~30 km thick

Jurassic rift fabric

Subsequent subsidence allowed thick sediment accumulation.

<u>Key features:</u> EM: Ellsworth Mountains EWM: Ellsworth-Whitmore block PSZ: Pagano Shear Zone RFIS: Ronne Filchner Ice Shelf, overlies *Southern Weddell magmatic province*

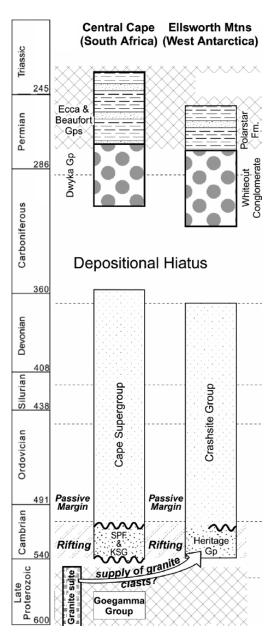
Jordan et al. 2017, doi: 10.1016/j.gr.2016.09.009

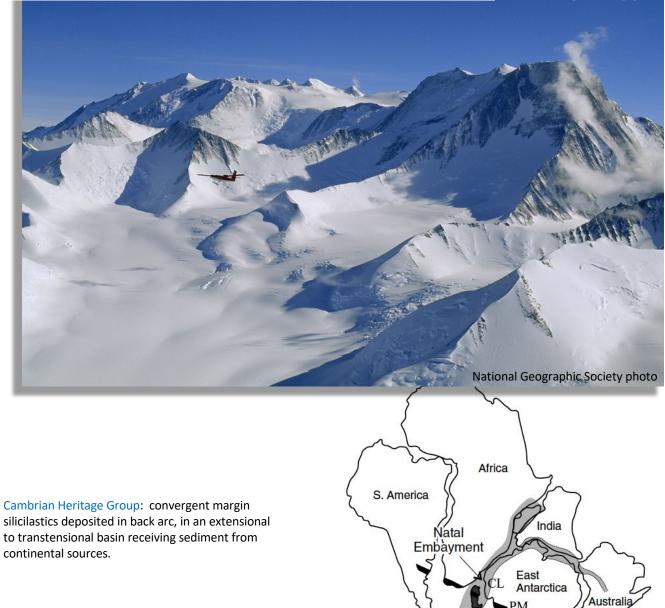




EWM Paleozoic sedimentary record, structural trend, and paleomagnetic data

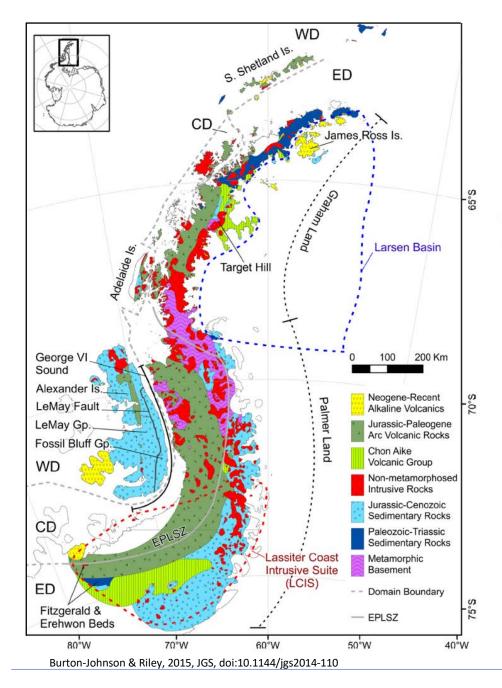
Mt. Vinson



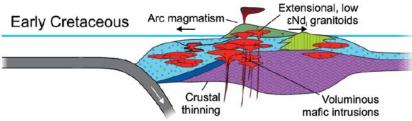


Randall & Mac Niocaill 2004 GJI

EWM



Main phase of continental margin arc magmatism in AP and TI : Cretaceous

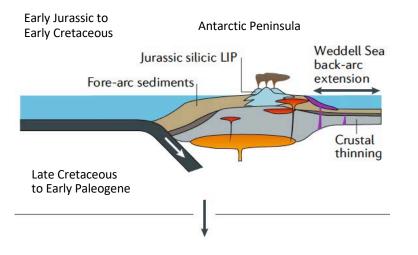


Debated geology: Amalgamation of far-travelled terranes that accreted to the Gondwana margin? Or not far travelled? (Patagonia is coherent)

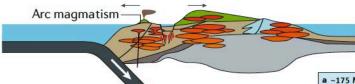
The geological record includes well-preserved forearc, back-arc and magmatic successions of a continental margin arc. The accretionary continental margin has been long-lived in this sector (Ordovician to present), with especially high magma addition rates during the Cretaceous and Permian. The Permian granite magmatism, ca. 280-250 Ma, affected the AP, north Patagonian Massif and Choiyoi province.

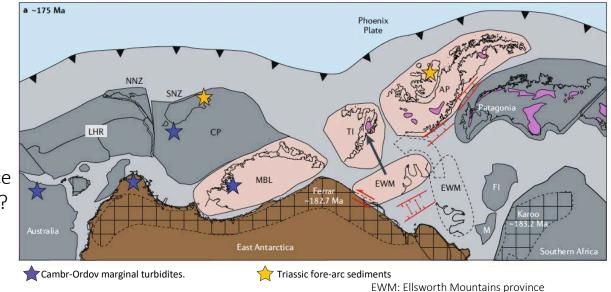


Ongoing subduction and magmatism in Antarctic Peninsula, with only limited extension across the arc.



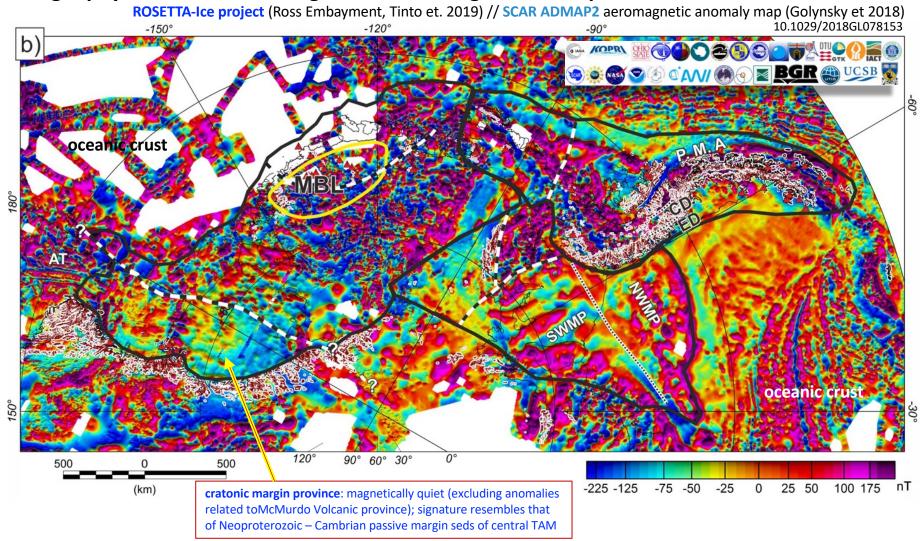






In the **back-arc region**: tectonism in southern Weddell magmatic province

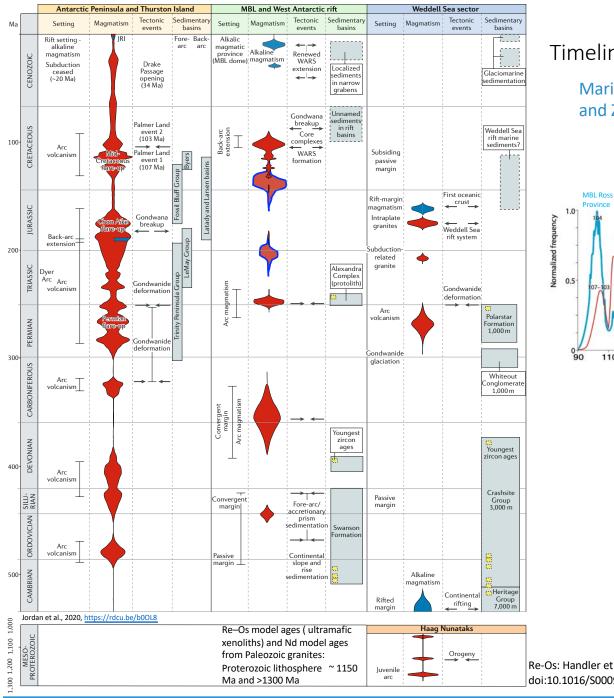
- wrench and ~30° tectonic rotation?
- >90° tectonic rotation?



New geophysical datasets – subglacial extent of geotectonic provinces

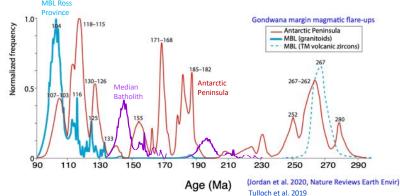
Marie Byrd Land (MBL) and West Antarctic rift system (WARS) are characterized by broad areas of crust 20–25 km thick, a product of Cretaceous to Cenozoic extension. The disordered high-amplitude magnetic patterns are attributed to wide-spread plutonic centers and volcanoes. Pacific-margin Magnetic Anomaly (PMA), ~2,600-km-long: attributed to well-preserved magmatic arc. There is detailed structure within the magnetic anomalies, attributable to periods of contraction and crustal thickening.

ONLINE



Timeline of tectonic – magmatic events

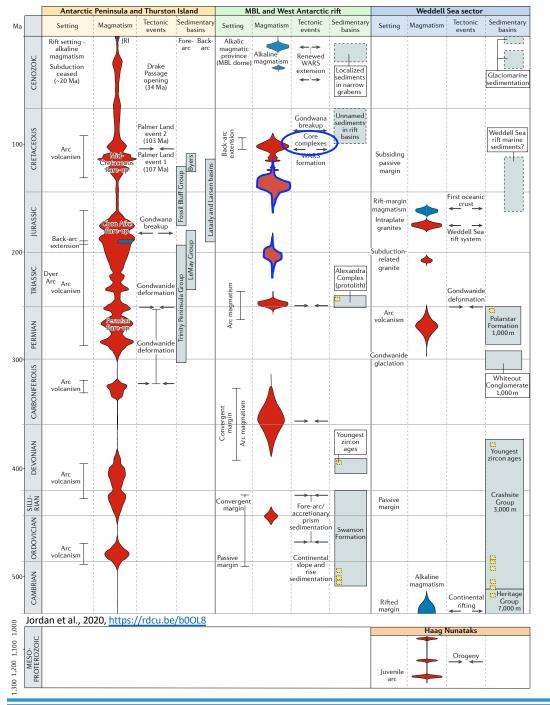
Marie Byrd Land, West Antarctic rift system and Zealandia *



* **Zealandia**: New Zealand plus submerged continental crust (Mortimer, 2017, doi: 10.1130/GSATG321A.1

Re-Os: Handler et 2003, doi:10.1016/S0009-2541(02)00410-2





Timeline of tectonic – magmatic events

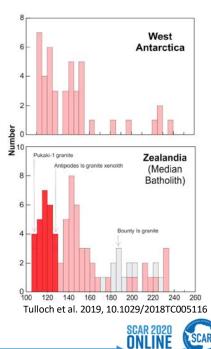
Marie Byrd Land + Zealandia



Base map: reconstruction of Tulloch et al. 2006

Broad margin with Mesozoic magmatic arc traversing Zealandia in the outboard position

Flareup with short voluminous magmatism is recorded in Median Batholith, Zealandia



MBL magmatic flare-up:

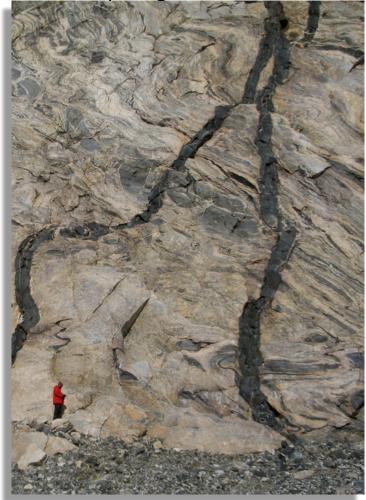
Bowman Peak – Byrd Coast Granite 102.9 ±0.9 Ma



Granulite facies metamorphism, crustal melting, core complexes formed and cogenetic granites emplaced. Dramatically rapid cooling of footwall rocks and granites.

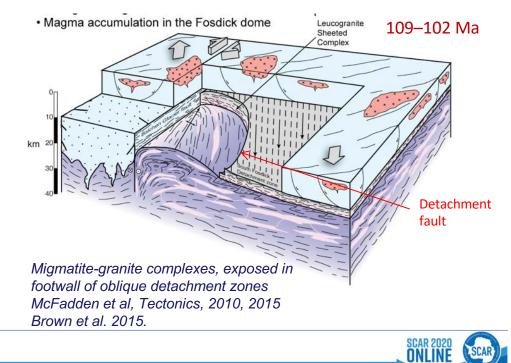


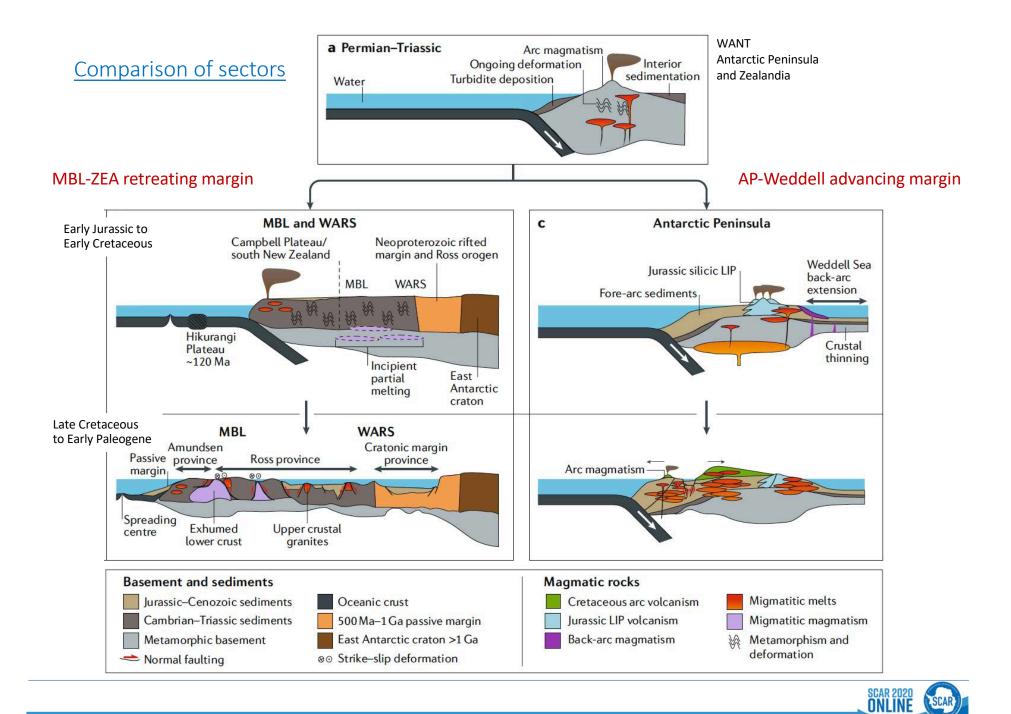
MBL flare-up: migmatization and mafic dyking



Melting and weakening of of mid-crust, migration to form granites. Folding and flow of melt-rich middle crust. Rapid exhumation upon detachment zones and wide extension.







What of the active volcanism ?

undergone repeated tectonic reworking with development of enduring crustal scale structures; underlain by anomalous mantle.

Pictured here: Basanite-trachyte flows, ~1.4 to 1.1. Ma, containing ultramafic xenoliths of varied composition *See Gaffney & Siddoway, ISAES 2007*

Marie Byrd Land volcanic province, 13 Ma to present (plus few older centers).



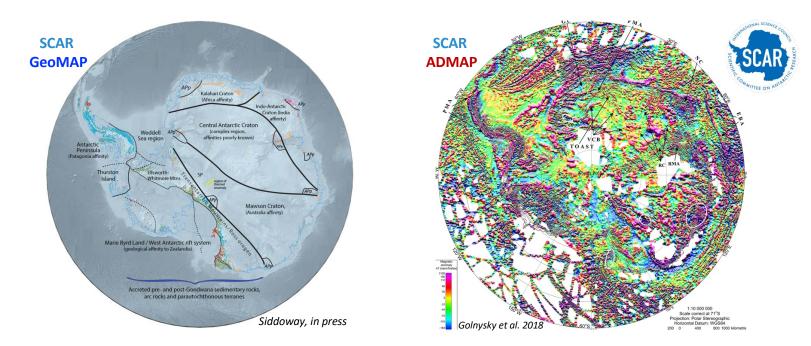
Credit: Tim Ivanic & Chris Yakumchuk

PHOTO: Demas Bluff, Fosdick Mountains – lavas containing deep sourced, diverse assemblage of ultramafic xenoliths



Conclusions (Part I)

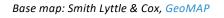
Geological analysis of the exposed rocks and geophysical observations together provide the means to unveil more of the region's complex history.



<u>The updated framework</u> presented here is a product of this joint approach. Key elements are:

- 1) the broader subglacial extent of WANT tectonic provinces (terranes)
- 2) varied tectonic-magmatic processes led to creation and stabilization WANT-Zealandia continental crust along an accretionary margin, tectonically active since 500 Ma
- 3) thin crust and extensive magmatism are a sign of repeated tectonic reworking that created a) distinctions between terranes and b) a common geotectonic history.
- 4) available for testing and refinement, and evaluation in the context of new research initiatives, including the new Lithosphere of East Antarctica program.

First meeting 12 Aug. 21:00 UTC



Acknowledgements

















Updated tectonic framework of West Antarctica

Siddoway, Jordan, Riley, Tinto & Tankersley

Place in the Presentation: beginning of Part II

Part I, completed

I. Geological/geophysical/ geochronological evidence of a shared history for West Antarctica tectonic provinces along a complex accretionary margin

II. Geophysical remote sensing reveals the subglacial extent of geological provinces; new tectonic scenarios emerge.

- Ellsworth-Whitmore province (Jordan et al. 2017)
- Antarctic Peninsula (Burton-Johnson & Riley, 2015)
- Ross Embayment cratonic margin province (Tinto et al. 2019)

Part II, ahead

- III. The inboard limit of West Antarctica in Ross Embayment
- Defined by crustal geology or physiography?
- Transantarctic Mountains or Central High as boundary of cratonic margin province?

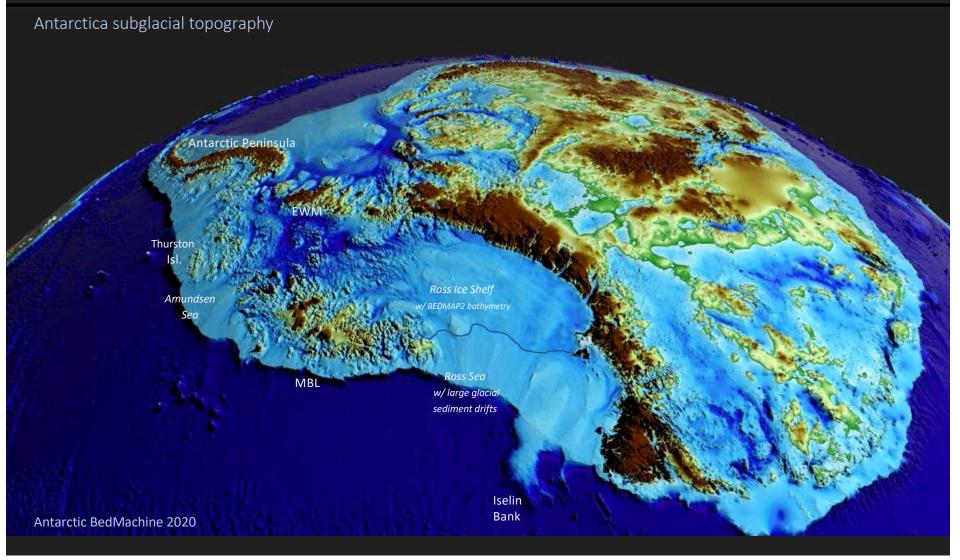
IV. Utility of updated crustal geology and tectonic framework, and hypothesis testing Relevant to many areas of contemporary research: paleotopography, glacial processes, subglacial hydrology, heat flux, glacial isostatic response



The inboard limit of West Antarctica in Ross Embayment

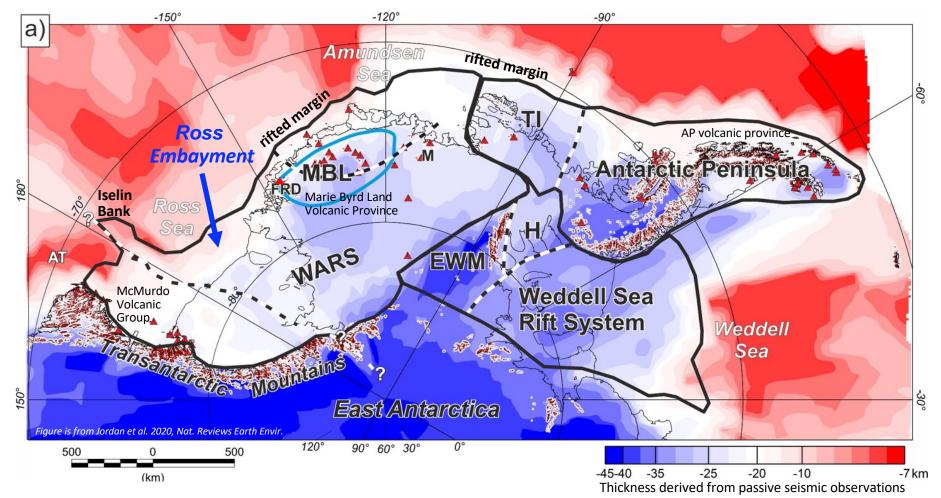
- Defined by crustal geology or physiography?
- Transantarctic Mountains or Central High as boundary of cratonic margin province?

Transantarctic Mountains form the physiographic boundary of East Antarctica.



Where is the inboard limit of West Antarctica within the Ross Embayment ?

Are there cues from the crustal thickness map and the distribution of Cenozoic volcanic provinces ?



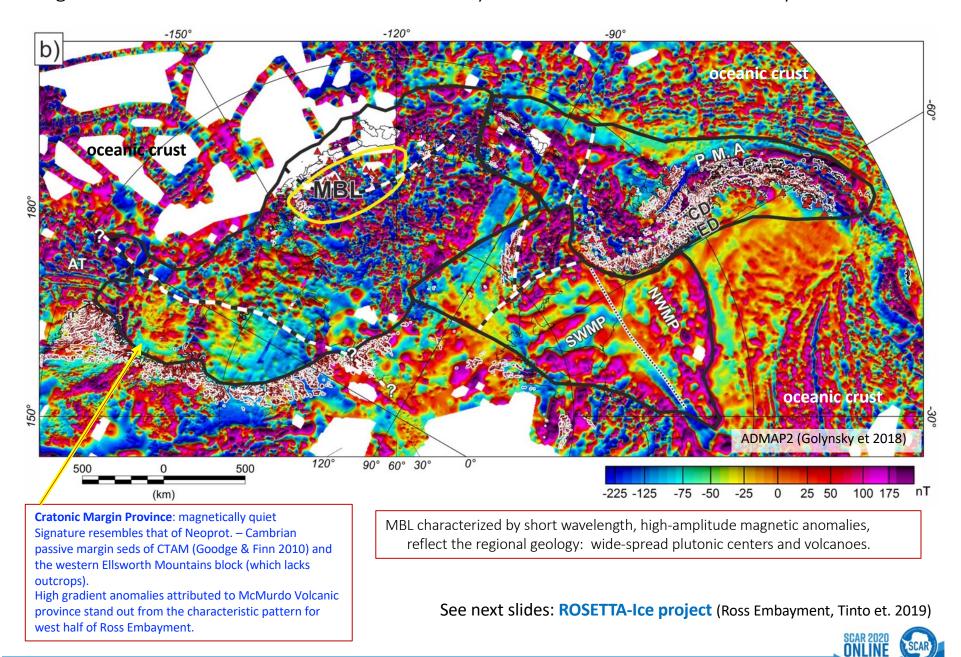
Major volcanoes formed 13 Ma to present, with isolated instances of Oligocene or earlier activity.

Heat flow measurements: 83 to 126 mW/m², some localities >250 mW/m²

Crustal thickness: EANT, 35–45-km. Ross Embayment, 15-28 km, and WANT, 22-32 km (Chaput et al. 2014).

A linear 'ridge' of thicker crust, ~28 km, is apparent along the midline of the Ross Embayment. It trends ~ N-S, parallel to the trend of the Iselin Bank 'promontory' upon the rifted margin. There are no known centers of volcanism along the ridge nor the bank.

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Magnetic character of MBL+ eastern Ross Embayment vs. western Ross Embayment

Higher resolution data for southern Ross Embayment collected by

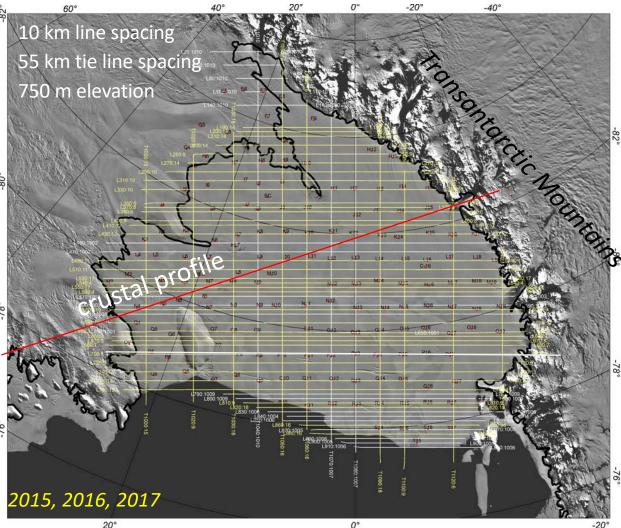
ROSETTA-Ice Project: Ross Ocean and Shelf Environment, and Tectonic setting Through Aerogeophysical surveys

Airborne survey 2015-2017 Ross Ice Shelf : 476,000 km² area Ice sheet 400 to >1000 m thick

Scintrex Cesium CS-3 magnetometer DGS gravity meter GNS gravity meter Shallow ice sounding radar Depth sounding radar RIEGL VQ-580 Swath scanning laser IMPERX Bobcat 29MP visible camera SOFRADIR IRE640L infrared camera Heitronics KT15-IIP Pyrometer



Two lines parallel to Operation IceBridge flight lines



Note the lines of longitude. North is down, in this figure.







Prominent contrast

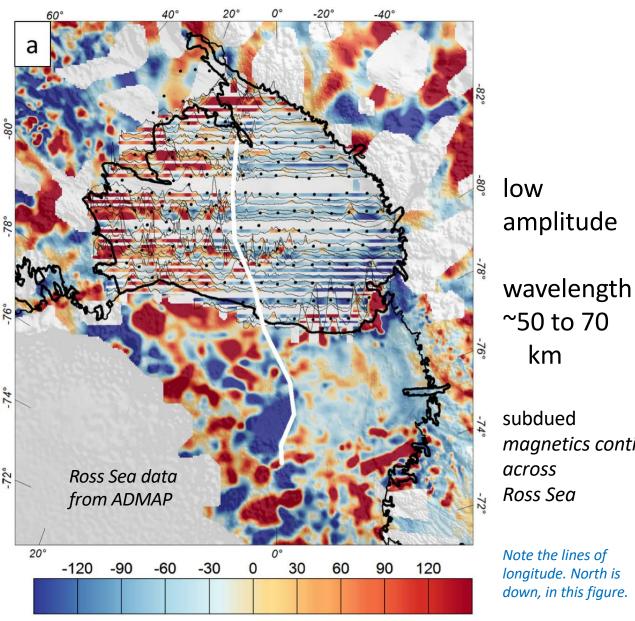
Magnetic Anomalies

high amplitude

wavelength ~20 km

Shallower depth to mag basement

High amplitude anomalies continue North across Ross Sea



Magnetic anomaly, nT

~50 to 70 magnetics continu

Note the lines of longitude. North is down, in this figure.



The Central High – prominent structure

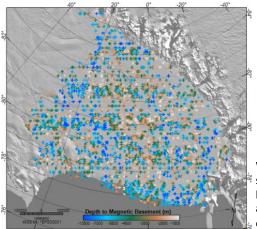
Southward continuation defined in this map of ROSETTA-Ice magnetic basement

Structures identified from seismic acoustic basement in Ross Sea, outlined in white. The major features are observed to continue southward beneath Ross Ice Shelf.

Rosetta-Ice Ross Ice Shelf data leveled to ANTOSTRAT acoustic basement, and joined to form a continuous surface.

Line pattern outlines the southward continuation of Ross Sea basement features.

Two ROSETTA-Ice flight lines coincide with Operation Ice Bridge flight lines (yellow dotted pattern).

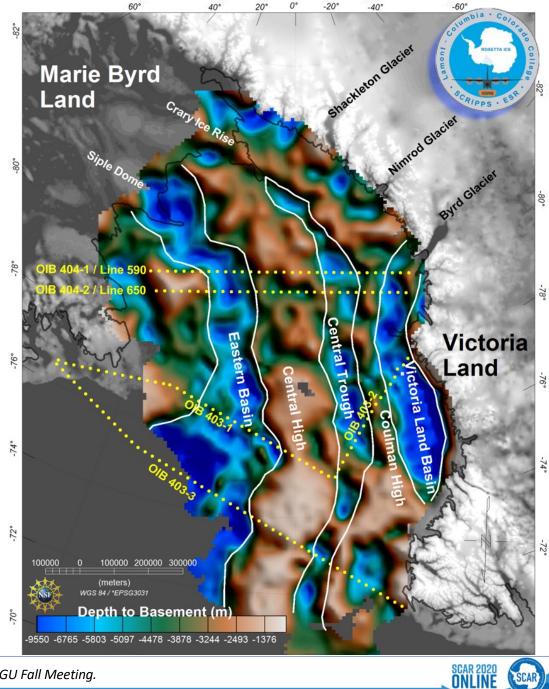


Werner deconvolution solutions with depth limits between sea level and 20km (Tankersley et 2018 AGU)

Work in progess:



M. Tankersley, Victoria Univ. Wellington, in progress; and 2018 AGU Fall Meeting.



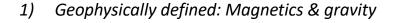
Our RIS crustal profile – portrays extended "EANT-type" crust beneath Ross Embayment Ice sheet 400 to >1000 m thick, upon ocean

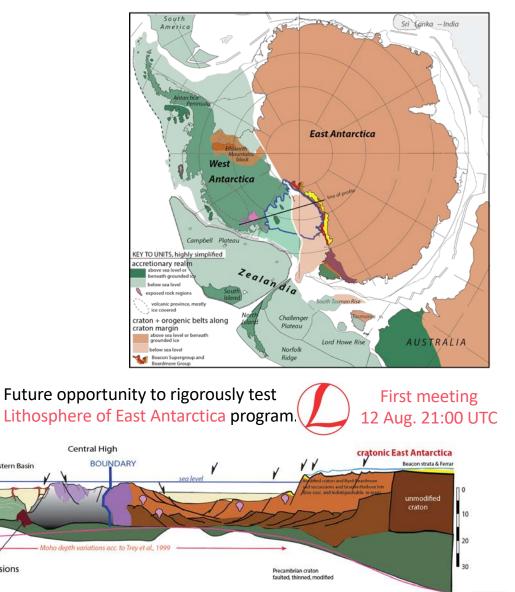
Roosevelt

Island

Eastern Basin

mafic intrusions





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Tinto et al. 2019, Nature Geoscience, doi 10.1038/s41561-019-0370-2

l velocity co

200 km

KE7 Plateau

of 0.1 to 0.3 km/s

gneiss domes

Line of profile, on map

Marie Byrd Land

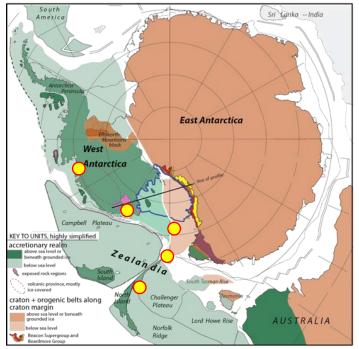
Fosdick Mtns @ V @ S Ford Rgs

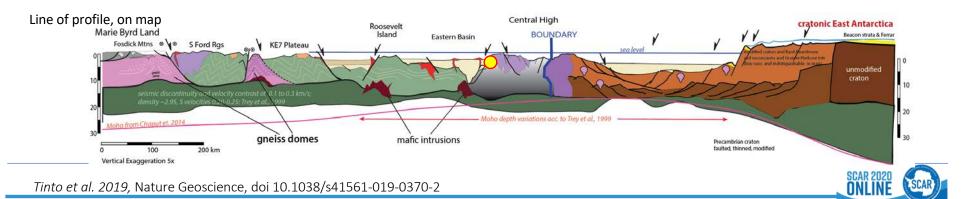
n Chaput et, 2014

100 Vertical Exaggeration 5x

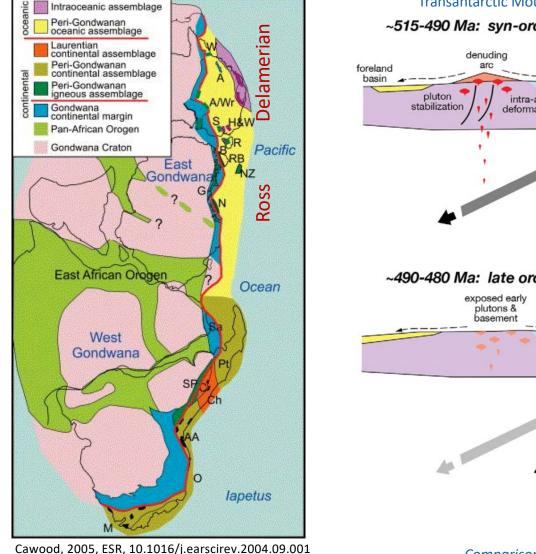
Our RIS crustal profile – portrays extended "EANT-type" crust beneath Ross Embayment Reasons why the interpretation is geologically valid

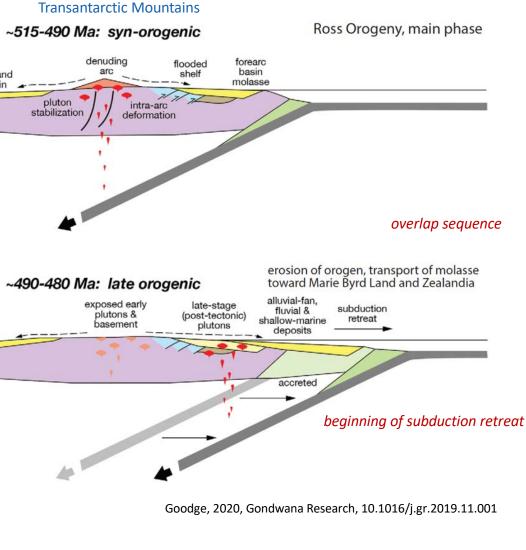
- 1) Geophysical characteristics: Magnetics & gravity (Tinto et al., 2019)
- 2) Zealandia (once contiguous with MBL, see map) restores to MBL, with edge of Campbell Plateau 'tipping out' at Iselin Bank. *Logically, the RIS "MBL-side" is Campbell Plateau-type crust, and the TAM-side is not Campbell Plateau type-crust.*
- 3) Continental extension is achieved through displacements on arrays of faults throughout the extended region, each of which accommodate a portion of the deformation. (not an individual fault that abruptly truncates the crustal geology see next slide)
- The TAM do not form a 'hard' boundary: Ross Orogen-type (~TAM)
 plutonic and metamorphic rocks exist outboard of the TAM (Ross Embayment, MBL, and ZEA Western Province)
 - overlap sequence of Cambrian-Ordov. siliciclastic sediment links CTAM and MBL
- 5) As recently as Oligocene, the Central High formed the prominent physiographic edge of Antarctica; see slide (not the TAM, as today)





Cambrian – Ordovician Ross – Delamerian Orogen





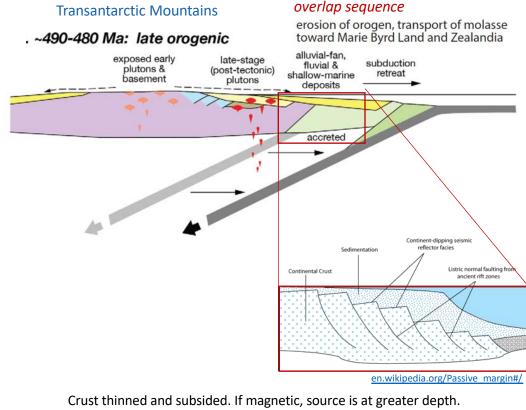
Comparisons of detrital zircon U-Pb data for CTAM – Pensacola versus Ford Ranges (MBL) suggest common sources from some units



Cambrian – Ordovician Ross – Delamerian Orogen



Cawood, 2005, ESR, 10.1016/j.earscirev.2004.09.001



Crust thinned and subsided. If magnetic, source is at greater depth. Sediment cover, thicker: continentally derived, non/low magnetism due to subaerial weathering and transport

Future opportunity to test location: Lithosphere of East Antarctica program.

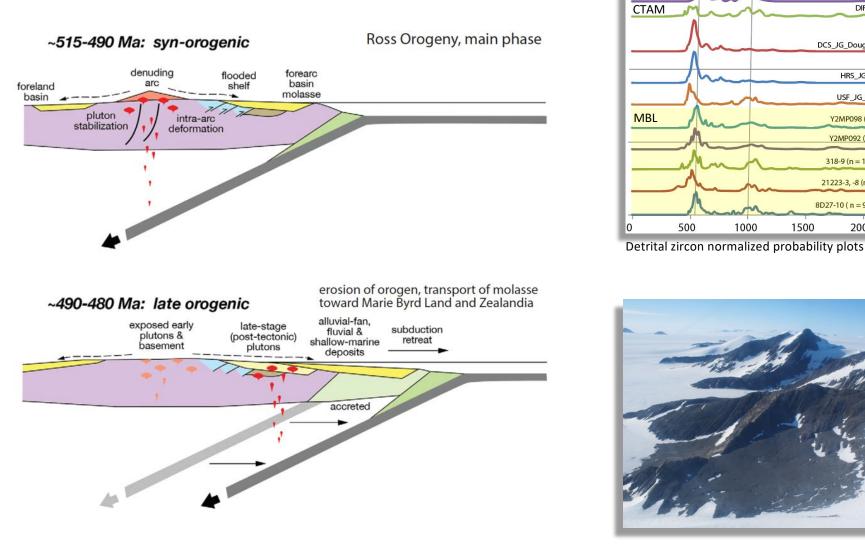


Comparisons of detrital zircon U-Pb data for CTAM – Pensacola versus Ford Ranges (MBL) suggest common sources from some units

Upper profile: Goodge, 2020, Gondw.Res., 10.1016/j.gr.2019.11.001



Cambrian – Ordovician Ross Orogen in Transantarctic Mountains



MBL data: Yakymchuk et al. 2015. TAM data: Goodge et al. 2004

Pensacola



PNA_JG_Patuxent (n = 65)

NHR_JG_Hannah (n = 72) DIF_JG_Byrd (n = 65)

DCS_JG_Douglas Congl. (n = 71)

HRS_JG_Starshot (n = 71)

USF_JG_Starshot (n = 101)

Y2MP098 (n = 91), Mt Passel

Y2MP092 (n = 89), Mt Dolber 318-9 (n = 175), Mitchell Peak 21223-3, -8 (n = 90), Scott Ntks 8D27-10 (n = 91), Mt Woodward

2000

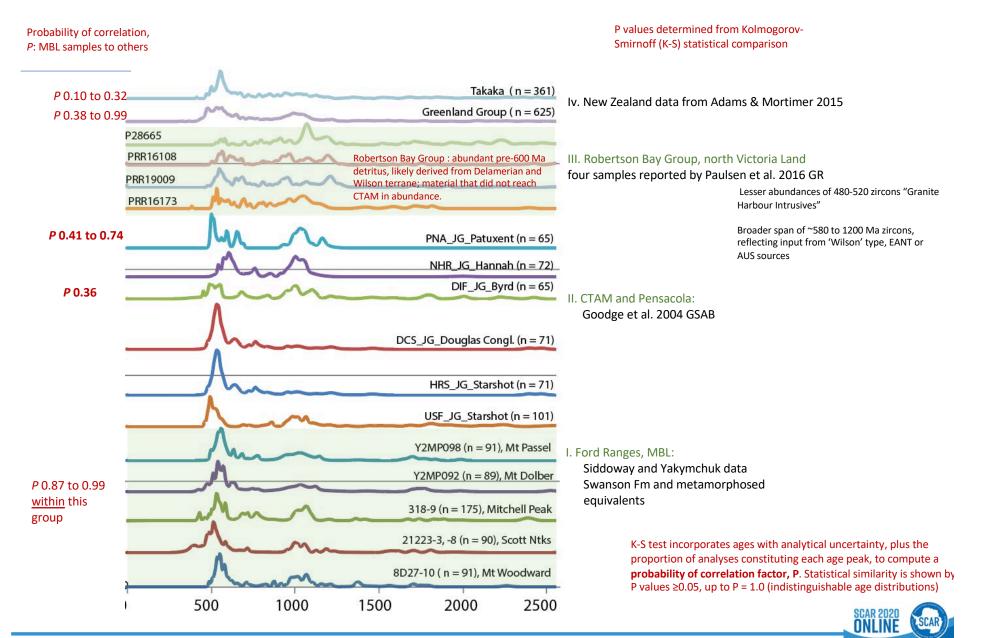
2500

1500

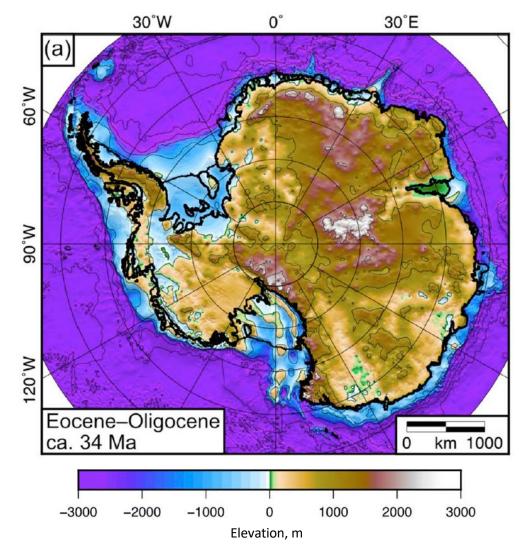
From J. Goodge, Gondwana Research 80 (2020), 10.1016/j.gr.2019.11.001

Latest Neoproterozoic to Cambrian orogenic sediment (molasse) shed from Ross Orogen

Normalized probability plots of detrital zircon U-Pb age data



5) As recently as Oligocene, the Central High formed the formed the prominent physiographic edge of Antarctica



Paxman et al. 2019, P³, doi:10.1016/j.palaeo.2019.109346

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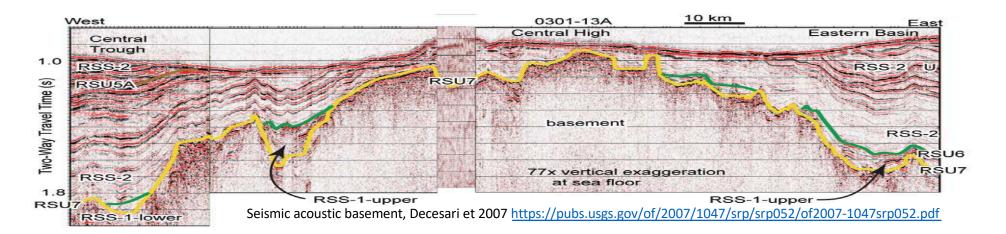
Bed corrected for ice load, thermal contraction, sedimentation, erosion, and horizontal motion

What defines the tectonic boundary and what is the geological character of the Central High?

HYPOTHESES to be tested:

Siddoway, Tankersley & Tinto, in progress

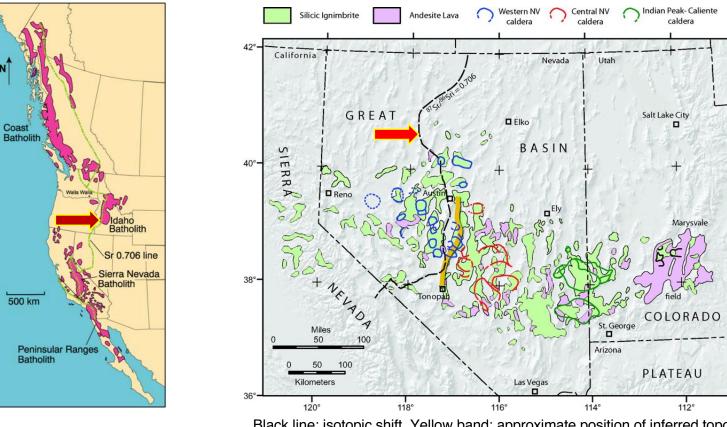
- Faulted thinned edge of continental crust of EANT character -- outer limit of evolved lithosphere. Analogy: USA Great Basin, and ⁸⁷Sr/⁸⁶Sr line. (" the .706 line"). *Possible that all major normal faults dip the same way*, toward MBL.
- 2) Fault-dissected continental block (ribbon) within a 'failed' volcanic passive margin . "Passive margin" in this case means the region of broad intracontinental extension formed in the Cretaceous. *Opposite dip directions in major crustal detachment faults on either side of the continental block.* Analogy: Voring margin, Norway
- 3) Thick crust bordering ancestral transform boundary that was active during Gondwana breakup, circa 85 Ma [see Lamb et al. 2016 Alpine Fault origins]
- 4) Microcontinent detached from continent (TAM) [Rey & Muller 2009 model]. Analogy: western USA terranes, NVL terranes
- 5) Linear plutonic belt (intermed to felsic magmatic arc) competent/coherent because is ~isotropic (previously unfaulted and not internally deformed) *during or after Ross Orogen.* Analogy: Andean type magmatic arc



Five Hypotheses to be tested: TWO viable analogs illustrated here

Hypothesis 1: North America – Great Basin as WARS analog

the 0.706 line : ⁸⁷Sr/⁸⁶Sr-delineated western edge of North American craton (Laurentia) within faulted, thinned, ancient passive margin of Rodinia



Black line: isotopic shift. Yellow band: approximate position of inferred topographic barrier on the western lip of the Nevadiplano "middle Cenozoic." Red and blue: calderas localized along the 0.706 line



Five Hypotheses to be tested: TWO viable analogs illustrated here

Hypothesis 2: Fault-dissected continental block (ribbon)

North Atlantic continent-ocean boundary and 'failed' volcanic passive margin as analog

Norway: Voring marginal high (e.g. *Gernignon et al ESR 2020*) 6704/12-1 Vorina SE Marginal High N. Gjallar Ridge Vigrid Syncline Fles Fault Complex Traena Basin 0 2 And and a state of the state of 6 Depth (km) 8 and a second second second 10 Constantine and the second T Reflection Presentation and presson Vp ~ 6.25km/s Vp ~ 7.1 km/s

Ross Embayment – 'failed volcanic passive margin ?

Evidence of structurally controlled mafic bodies: sills and steep, narrow zones

Steep zones have trends <u>orthogonal</u> to rifted margin (subparallel to TAM)

<u>Geophysics</u>: **ANTOSTRAT** and **ACRUP** marine geophysics (e.g. Busetti et al. 1999; Della Vedova et al. 1994- 98, Davey et al. 1995, Cooper et al. 1991, 1997)

ROSETTA-Ice (Tinto et al. 2018)

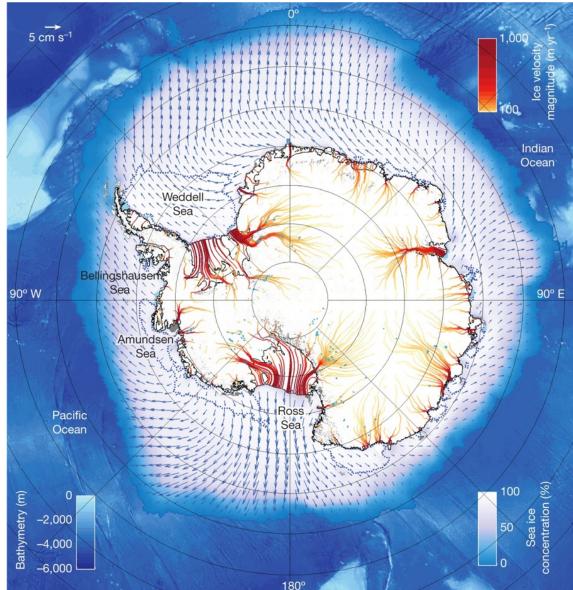


Concluding thoughts: Updated crustal geology and tectonic framework

Bedrock geology and tectonic context are directly relevant to contemporary research in many areas:

- modeling the Cryosphere
- Geosphere Ocean system
- paleotopography for past climate reconstructions
- bedform influences upon glacial processes
- subglacial hydrology
- heat flux
- glacial isostatic response

Figure: Ice sheet motion flowlines superimposed on MODIS mosaic of Antarctica Sea ice motion vectors



Shepherd, Fricker & Farrell, 2018, https://www.nature.com/articles/s41586-018-0171-6/figures/1



Concluding thoughts: Updated crustal geology and tectonic framework

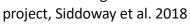
Synergistic links: International Ocean Discovery Program circum-Antarctica drilling

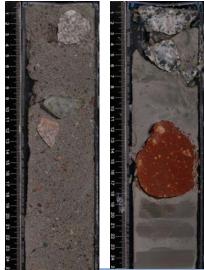
Opportunity to further characterize subglacial bedrock and WANT geological evolution using iceberg-rafted detritus, including clasts of rock, derived from WANT and EANT



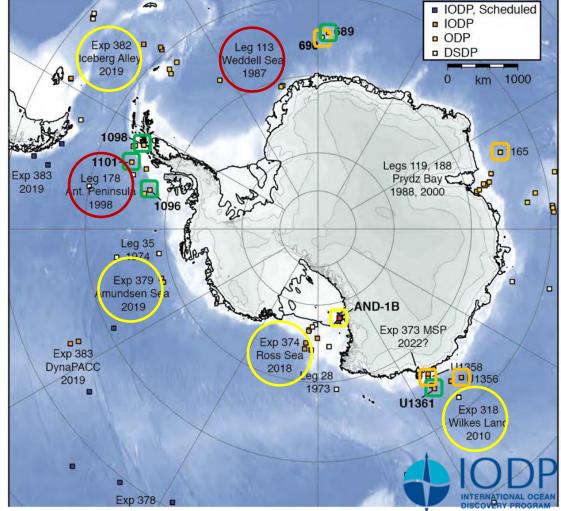
IODP379 Shipboard Party

Geo/thermochronology data (multi-chronometry upon individual grains, to further refine tectonomagmatic development and landscape evolution of WANT e.g. NSF ICI-Hot





Igneous clasts within cores, IODP379



Map: IODP, Kulhanek & Williams 2019



Acknowledgements















